

3.13 EXAMPLE EXERCISE ON ELASTICITY PLUS WORKED ANSWER

Question

- List the principal properties of isotropic and anisotropic elastic materials, and give examples of such materials.
- Define the quantities "Young's Modulus" and "Poisson's Ratio" used to specify the elastic behaviour of isotropic materials.
- Write down a set of equations giving the 3-D strain response of an isotropic elastic solid subjected to a state of stress defined by the principal stresses $\sigma_1, \sigma_2, \sigma_3$.
- A block of isotropic quartzite (Young's Modulus $E = 3 \times 10^{10}$ Pa, Poisson's ratio $\nu = 0.3$) is subjected to a state of stress defined by the tensor

$$\sigma_{ij} = \begin{bmatrix} 90 & 0 & 0 \\ 0 & 60 & 0 \\ 0 & 0 & 30 \end{bmatrix} \text{ MPa}$$

where compressive stress is taken as positive.

Calculate the resulting state of elastic strain, writing your answer in tensor notation.

Worked Answer

- The principal properties of isotropic elastic materials are as follows:
 - The elastic constants (such as Young's Modulus E or shear modulus G) are independent of direction.
 - The principal stresses are parallel to the principal strains
 - Under elevated hydrostatic pressure P , a sphere of elastically isotropic material compresses to a smaller sphere
 - A bar of isotropic material can be stretched uniaxially without shearing, and can be bent without twisting

In contrast, elastically anisotropic materials possess the following principal properties:

- The elastic constants depend on orientation and are related to material symmetry.
- The principal stresses are generally not parallel to the principal strains.
- Under elevated hydrostatic pressure, a sphere becomes an ellipsoid.
- A bar of anisotropic material will generally shear when stretched and twist when bent.

Examples are given below.

- Elastically isotropic materials: polycrystalline solids with no crystallographic preferred orientation (e.g. statically recrystallized marble or metal), structureless rocks (e.g. massive quartzite or limestone).
 - Elastically anisotropic materials: polycrystalline solids with a well-developed crystallographic preferred orientation (e.g. strongly deformed rocks or metals), foliated or layered rocks (e.g. slates or laminated sediments) anisotropically cracked materials, all single crystals.
- b) Young's modulus (E) and Poisson's ratio (ν) are defined for uniaxial loading of an isotropic elastic material as follows:-

$$E = \sigma_1/\varepsilon_1 = d\sigma_1/d\varepsilon_1 = \text{slope of the linear stress } (\sigma_1) \text{ versus axial strain } (\varepsilon_1) \text{ curve obtained in uniaxial loading.}$$

$$\nu = -\varepsilon_1/\varepsilon_2 \quad \text{where } \varepsilon_2 = \varepsilon_3 \text{ is the lateral strain obtained in uniaxial loading.}$$

c) The 3-D strain response of an elastically isotropic material is written

$$\varepsilon_1 = [\nu/E](\sigma_1/\nu - \sigma_2 - \sigma_3)$$

$$\varepsilon_2 = [\nu/E](\sigma_2/\nu - \sigma_1 - \sigma_3)$$

$$\varepsilon_3 = [\nu/E](\sigma_3/\nu - \sigma_1 - \sigma_2)$$

where the ε_i are the principal strains and the σ_i are the principal stresses.

d) For the block of quartzite, we have (given)

$$E = 3 \times 10^{10} \text{Pa}$$

$$\nu = 0.3$$

Furthermore, from the given values of σ_{ij} (and noting that the shear stress components are all zero) we can directly write the principal stresses as

$$\sigma_1 = \sigma_{11} = 90 \text{ MPa}$$

$$\sigma_2 = \sigma_{22} = 60 \text{ MPa}$$

$$\sigma_3 = \sigma_{33} = 30 \text{ MPa}$$

Using the 3-D relations given under (c) hence leads to the result

$$\varepsilon_1 = \frac{0.3}{3 \cdot 10^{10}} \cdot \left(\frac{90}{0.3} - 60 - 30 \right) \cdot 10^6 = 2.1 \cdot 10^{-3}$$

$$\varepsilon_2 = \frac{0.3}{3 \cdot 10^{10}} \cdot \left(\frac{60}{0.3} - 90 - 30 \right) \cdot 10^6 = 0.8 \cdot 10^{-3}$$

$$\varepsilon_3 = \frac{0.3}{3 \cdot 10^{10}} \cdot \left(\frac{30}{0.3} - 90 - 60 \right) \cdot 10^6 = -0.5 \cdot 10^{-3}$$

which in tensor notation can be written

$$\varepsilon_{ij} = \begin{bmatrix} 2.1 & 0 & 0 \\ 0 & 0.8 & 0 \\ 0 & 0 & -0.5 \end{bmatrix} \cdot 10^{-3}$$

4.11 MODEL ANSWER TO A PREVIOUS EXAM QUESTION

How to SCORE in your exam!

QUESTION

- a) Explain what is meant by the term "failure criterion".
- b) Write down the Coulomb criterion for shear failure of isotropic rock when pore fluid is present at pressure P_f . Identify all terms!
- c) Write down an expression giving the failure plane orientation implied by the Coulomb criterion.
- d) The earthquake of 1992 on the Peelrand fault near Roermond occurred at a depth of $h \approx 23$ km. Noting that the Peelrand fault is a normal fault bounding a graben structure, and assuming that the fault is characterized by a cohesive shear strength of 1 MPa and a coefficient of internal friction of $1/2$, estimate the horizontal stress (σ_3) associated with the triggering of the earthquake. (These data are based on experimental data for long-inactive, healed faults).

HINT: Assume that σ_1 is subvertical and equal to the lithostatic overburden pressure ($p = \rho gh$) at $h = 23$ km depth, and that the pore fluid pressure at that depth is $0.8 p$ as frequently observed in boreholes. Take the density (ρ) of the overburden to be 2500 kg/m^3 and $g = 10 \text{ ms}^{-2}$.

- e) Comment on the depth at which the Roermond earthquake occurred.

ANSWER

- a) A failure criterion is a mathematical relation specifying the states of stress (σ_{ij}) at which brittle failure of a material occurs.
- b) The Coulomb criterion for shear failure of wet rock is written:

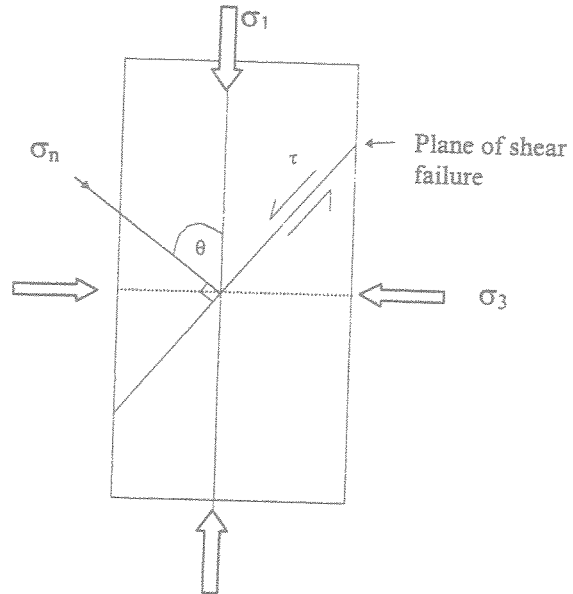
$$\tau = S_0 + \mu (\sigma_n - P_f)$$

where τ is the shear stress on the failure plane
 S_0 is the cohesive shear strength of the rock
 μ is the so-called coefficient of internal friction

σ_n is the normal stress acting across the failure plane
 P_f is the pore fluid pressure

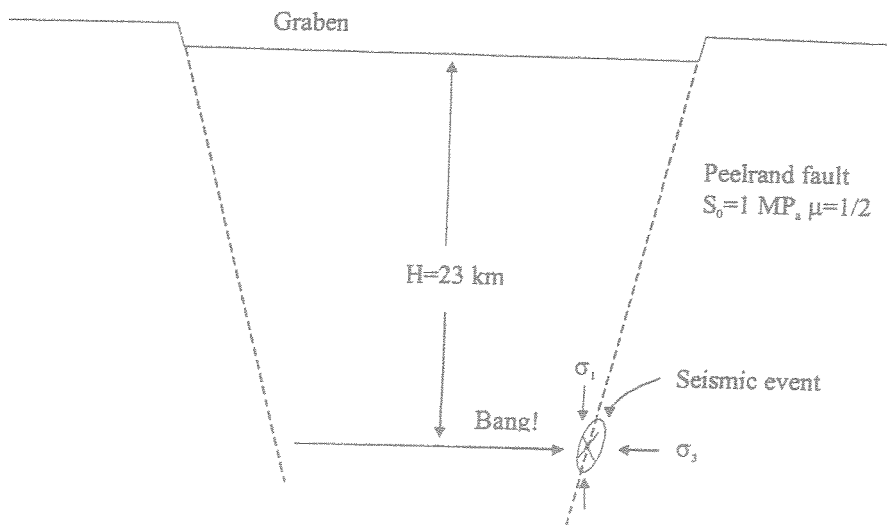
c) The failure plane orientation implied by the Coulomb criterion is given

$$\tan 2\theta = \frac{1}{\mu}$$



where 2θ lies in the range $90^\circ \leq 2\theta \leq 180^\circ$
 and θ is defined as shown above right.

d) Problem: Determine σ_3 associated with the triggering of seismic slip, in the following situation:-



From the given information, we have: $S_0 = 1 \text{ MPa}$ and $\mu = 1/2$

Hence we can write, for the fault, the Coulomb criterion

$$\tau = 1 + (1/2)(\sigma_n - P_f) \quad (\text{MPa}) \quad (1)$$

In addition, from the equations for 2-D stress we can write the following equations for the shear stress (τ) and normal stress (σ_n) on the fault plane

$$\tau = [1/2](\sigma_1 - \sigma_3) \sin 2\theta \quad (2)$$

$$\sigma_n = [1/2](\sigma_1 + \sigma_3) + [1/2](\sigma_1 - \sigma_3) \cos 2\theta \quad (3)$$

At failure, we must therefore have (from 1, 2, 3)

$$\frac{1}{2}(\sigma_1 - \sigma_3) \sin 2\theta = 1 + \frac{1}{2} \left\{ \frac{\sigma_1 + \sigma_3}{2} + \frac{(\sigma_1 - \sigma_3)}{2} \cos 2\theta - P_f \right\} \quad (4)$$

Now, from the given information

$$\sigma_1 (\text{vertical}) = \rho gh = (2500)(10)(23000) = 575 \text{ MPa}$$

$$P_f = (0.8)\rho gh = 460 \text{ MPa}$$

and from the Coulomb criterion (eqn. 1)

$$\tan 2\theta = -\frac{1}{\mu} = -2 \quad (90^\circ \leq 2\theta \leq 180^\circ)$$

or $2\theta = 116.6^\circ$ (students, be careful!! Note range of 2θ specified above)

Putting these values for σ_1 , P_f and 2θ into (4) now yields

$$(575 - \sigma_3)(0.89) = 2 + \frac{575 + \sigma_3}{2} + \left(\frac{575 - \sigma_3}{2} \right)(-0.45) - 460$$

which on rearranging gives

$$\sigma_3 = 1623.25/3.23 = 503 \text{ MPa}$$

Hence the horizontal stress associated with triggering seismic slip on the fault is

$$\sigma_3 = 503 \text{ MPa} \approx 500 \text{ MPa}$$

- e) The Roermund earthquake occurred well below the depth of 10-15 km at which the brittle-ductile transition normally occurs. This could reflect either usually low values of σ_3 at depth (perhaps caused by an acceleration of crustal extension) or unusually high values of P_f . Both could cause brittle failure at depths where ductile behaviour is normally expected. Another possible explanation is an unusually low geothermal gradient, giving a deeper brittle-ductile transition. However, the geotherm is not unusually low in the Rhine-graben region.

4.12 EFFECT OF PORE FLUID PRESSURE ON ELASTIC BEHAVIOUR

4.12.1 Basic Theory

Before leaving the general topic of the behaviour of rock under upper crustal (elastic-brittle) conditions, let us consider the effects of pore fluid pressure (P_f) on elastic behaviour.

For the elastic response of isotropic rock in the absence of pore fluid pressure we have already had (taking compressive stresses and strains as positive)

$$2G \epsilon_{ij} = \sigma_{ij} - \frac{1}{3}(1 - 2G/3K) \sigma_{kk} \delta_{ij}$$

shear modulus G
 $\bar{\sigma} = \Delta P$
 bulk modulus K

$$\bar{\sigma} = \frac{1}{3}(\sigma_{11} + \sigma_{22} + \sigma_{33})$$

$$K = \frac{\bar{\sigma}}{\epsilon_v} = \frac{\Delta P}{\epsilon_v} \quad (1)$$

where σ_{ij} is the stress applied to the rock, ϵ_{ij} is the resulting strain, G is the shear modulus and K the bulk modulus of the *bulk dry rock*. If the rock is porous and pore fluid is introduced at pressure P_f , the normal components of (effective) stress supported by the solid framework will be reduced by an amount P_f . Thus, the mean (effective) normal stress will change by an amount $\Delta \sigma = -P_f$. This will produce an isotropic elastic volumetric expansion strain given (from the definition of K) as

$$\epsilon_v^- = \Delta \bar{\sigma} / K = -P_f / K = \epsilon_{11}^- + \epsilon_{22}^- + \epsilon_{33}^-$$

$$\epsilon_{ij}^- = -P_f \delta_{ij} / 3K \quad (\text{since } \epsilon_{11}^- = \epsilon_{22}^- = \epsilon_{33}^-)$$

$\Delta \bar{\sigma} = -P_f$
 volumetric strain ϵ_v^-

At the same time, P_f will cause an isotropic volumetric contraction of the grains (solid framework) of the rock given (from the definition of the bulk modulus of the grains, K_g)

$$\epsilon_v^+ = P_f / K_g = \epsilon_{11}^+ + \epsilon_{22}^+ + \epsilon_{33}^+$$

$$\epsilon_{ij}^+ = P_f \delta_{ij} / 3K_g \quad (\text{since } \epsilon_{11}^+ = \epsilon_{22}^+ = \epsilon_{33}^+)$$

$\Delta \bar{\sigma} = +P_f$
 volumetric strain ϵ_v^+

bulk modulus of grains